

# SYNERGETIC AEROSOL RETRIEVAL FROM ENVISAT

Thomas Holzer-Popp, Marion Schroedter-Homscheidt

*Deutsches Zentrum für Luft- und Raumfahrt e. V. (DLR), Deutsches Fernerkundungsdatenzentrum (DFD),  
Oberpfaffenhofen, 82234 Wessling, Germany, Email: thomas.holzer-popp@dlr.de*

## ABSTRACT

A synergetic aerosol retrieval method (SYNAER) to retrieve aerosol optical thickness and type has been developed and validated at DLR-DFD with GOME and ATSR-2, both onboard ERS-2 ([1] and [2]). This method is currently adapted to ENVISAT SCIAMACHY and AATSR and will also be applicable to METOP GOME-2 and AVHRR. Based on 14 months of data in 1997/1998 from ERS-2 over the MSG observation area a first climatological 5x5 degree data set has been derived, which describes annual average optical thickness values of major components (sulfate/nitrate, soot, dust, sea salt) of the atmospheric aerosol load. The paper summarizes the first preliminary results of the AO project SENECA: the ERS-2 climatology effort and the adaption to ENVISAT.

## 1. INTRODUCTION

Air pollution by solid and liquid particles suspended in the air, so-called aerosols is one of the major concerns in many countries on the globe. One focus of concern is related to aerosols from anthropogenic origin mainly by combustion processes (industry, vehicle transport, heating, biomass burning). In developed countries improved combustion and filtering methods have led to a general decrease of particle concentrations in total suspended matter but new concern arises from potential health impact of increasing numbers of smaller aerosols, so-called nano-particles in particular from diesel engines ([3]). [4] reports that 21 – 38 % of total excess deaths in the UK during the summer heatwave of 2003 are attributable to elevated ozone and particle concentrations. On the other hand developing countries suffer still from high total particle loads in the air. Furthermore, natural aerosols (mainly dust and sea salt) also contribute significantly to background and episodically severely increased particle concentrations. Dust can also act as carrier for long-range transport of diseases, e.g. from the Sahara to the Caribic or Western Europe ([5]), or even once around the globe ([6], as recently Chinese yellow sand was detected as far as the Swiss Alps. Also well known in principle are direct (by reflecting light back to space) and indirect (by acting as

cloud condensation nuclei) climate effect of aerosols, although large uncertainties exist in the exact values of the forcing. Recent publications (e.g. [7]) point out, that the absorption behaviour of particles (mainly soot and minerals) needs to be known to assess their total climate effect (strongly absorbing particles can regionally reverse the sign of the aerosol forcing from cooling to heating).

In the light of this overall picture, climate monitoring, long-term air pollution monitoring as well as short term forecast of pollution levels need to take into account intercontinental transport processes and the composition of the atmospheric particle load. Satellite observations of the total aerosol mass have experienced significant improvements in the last few years thanks to improved instrumentation and enhanced retrieval algorithms ([8]). Thus they offer the potential to regularly monitor the global aerosol distribution and by assimilating these measurements into chemistry-transport-models to enhance particle forecasts especially for episodic severe pollution events. Furthermore, satellite data can contribute to deduce background and long-range transport patterns of aerosols. And finally, with the emerging capacity to separate the total aerosol mass observed with optical observations into major components, a better understanding of the behaviour of different chemical and size classes is supported. As a first step towards this goal a satellite-based climatological data set of the optical thickness of major aerosol components over Europe / Africa is presented here which can be used as background aerosol distribution maps.

After a brief introduction in section 1 the retrieval method as well as the status of validation efforts against AERONET spectral sun photometer measurements and other satellite observations are described in section 2. Section 3 deals with the improvements in the aerosol model and section 4 describes the first preliminary results of the application of the synergetic method to the 14 month dataset over Europe and Africa. Section 5 gives an overview of the status of adapting SYNAER to ENVISAT AATSR and SCIAMACHY including first results. Sections 6 to 8 contain the summary, acknowledgements and references.

## 2. RETRIEVAL METHOD: OVERVIEW AND ITS VALIDATION STATUS

At the German Remote Sensing Data Center (DFD) the new aerosol retrieval method SYNAER (SYNergetic AERosol Retrieval) was developed ([1]) which delivers boundary layer aerosol optical thickness (BLAOT) and type over both land and ocean, the latter as percentage contribution of 6 representative components from the OPAC (Optical Parameters of Aerosols and Clouds, [9]) dataset to BLAOT. The high spatial resolution of the ATSR-2 instrument (Along Track Scanning Radiometer) permits accurate cloud detection, BLAOT calculation over automatically selected and characterized dark pixels and surface albedo correction for a set of 40 different pre-defined boundary layer aerosol mixtures. After spatial integration to the larger pixels of the spectrometer GOME (Global Ozone Monitoring Experiment) these parameters are used to simulate GOME spectra for the same set of different aerosol mixtures. A least square fit of these spectra to the measured spectrum delivers the correct BLAOT value and - if a uniqueness test is passed - the aerosol mixture. For humidity dependent components two models with 50% and 80% relative humidity have been included. SYNAER is currently under implementation for operational processing at the German Remote Sensing Data Center, which also serves as processing and archiving center (PAC) for all atmospheric ENVISAT data on behalf of ESA.

Accurate cloud detection is an important prerequisite for each aerosol retrieval. The well established APOLLO (AVHRR Processing Scheme Over CLOUD Land and Ocean; [10]) software was adapted to ATSR-2 data. The capability of retrieving cloud cover in boxes of 1 km<sup>2</sup> means a significant strength of SYNAER because it reduces the erroneous aerosol detection due to the presence of sub pixel clouds significantly. It even allows the correction of partly cloudy GOME pixels. First case studies distributed over the globe using ground based sun-photometer measurements of the spectral aerosol optical thickness from NASA's Aerosol Robotic Network (AERONET) at 14 locations and one airborne lidar measurement show a good agreement with errors less than 0.1 in six wavelengths (340 – 870 nm) which indicates correct assessment of the amount and type (namely the spectral dependence of extinction) of aerosol ([2]). This first groundbased validation comprised data from 4 continents in several climate zones (latitudes 17 South to 56 North) distributed over 2 years except the winter season with solar elevations ranging from 25 to 60 degrees. The cloud fraction

inside the GOME pixels ranged up to 35%. Furthermore, a comparison of monthly mean results from SYNAER and other satellite aerosol retrievals as well as AERONET stations over ocean ([11]) showed a satisfactory agreement with the other datasets.

## 3. IMPROVEMENTS IN THE AEROSOL MODEL

Basically the components (which summarize the optical features of all particles with similar optical behaviour) including their log-normal size distribution are taken from the OPAC database ([3]). Tab. 1 summarizes their relevant microphysical properties and derived (Mie-calculated) optical characteristics. However, more recent campaigns and AERONET data exploitations have been used to improve some specific items:

- the original soot component was split in 2 components for weakly absorbing biomass burning ([12]) and very strongly absorbing diesel (industrial) carbon ([13]). Furthermore, the real part of the refractive index was adjusted to [12] representing a mixture of Amazonian, South American cerrado, African savannah and boreal forest fires.
- in the absorption of mineral dust (transported and insoluble) 2 components were introduced to take the dust origin with different hematite content and consecutively absorption into account ([12] representing an average of Cape Verde, Saudi Arabia and Bahrain observations and [14] below 400 nm). In the case of desert outbreak (transported minerals) the lowest aerosol layer of 4-6 km was modeled as two distinct sub-layers (dust above background), as they occur in nature.

Tab. 2 shows the (updated) definition of the 40 mixtures used in the SYNAER retrieval method. Values in the table show the vertical profile, relative humidity in the boundary layer and the percentage contribution to the optical thickness at 550 nm of the respective components. Two groups of 20 mixtures, each are applied where either relative humidity or the hematite content of the mineral component (and consecutively the absorption) are altered. Alternative values are marked with grey boxes in tab. 2: For example, mixture number 1 has 50% relative humidity and mixture number 21 has 80% relative humidity; mixture number 2 has 5% insoluble (high hematite content) component, whereas mixture number 22 has 5% insoluble (low hematite content) contribution to the optical thickness at 550 nm.

Tab. 1: updated aerosol components (new components are highlighted)

Component	Species	Complex refract. Index at 550 nm	Mode radius [ $\mu\text{m}$ ]	Stand. Dev. of size distribution	Particle density [ $\text{g}/\text{cm}^3$ ]	Extinction coefficient for 1 particle per $\text{cm}^3$ at 550 nm [ $\text{km}^{-1}$ ]	Single scattering albedo At 550 nm	Literature source
WASO, rH = 70%	Sulfate/nitrate	1.53 – 0.0055 i	0.028	2.24	1.33	7.9 e-6	0.981	Hess, et al. 1998
INSO	Mineral dust, high hematite content	1.53 – 0.008 i	0.471	2.51	2.0	8.5 e-3	0.73	Hess, et al. 1998
<b>INSL</b>	<b>Mineral dust , low hematite content</b>	<b>1.53 – 0.0019 i</b>	<b>0.471</b>	<b>2.51</b>	<b>2.0</b>	<b>8.5 e-3</b>	<b>0.891</b>	<b>Dubovik, et al. 2002</b>
SSAM, rH = 70%	Sea salt, accumulation mode	1.49 – 0 i	0.378	2.03	1.2	3.14 e-3	1.0	Hess, et al. 1998
SSCM, rH = 70%	Sea salt, coarse mode	1.49 – 0 i	3.17	2.03	1.2	1.8 e-1	1.0	Hess, et al. 1998
<b>BISO</b>	<b>Biomass burning soot</b>	<b>1.63 - 0.036 i</b>	<b>0.0118</b>	<b>2.0</b>	<b>1.0</b>	<b>1.5 e-7</b>	<b>0.698</b>	<b>Dubovik, et al. 2002</b>
<b>DISO</b>	<b>Diesel soot</b>	<b>1.49 – 0.67 i</b>	<b>0.0118</b>	<b>2.0</b>	<b>1.0</b>	<b>7.8 e-7</b>	<b>0.125</b>	<b>Schnaiter, et al. 2003</b>
MITR	Transported minerals, high hematite content	1.53 – 0.0055 i	0.5	2.2	2.6	5.86 e-3	0.837	Hess, et al. 1998
<b>MILO</b>	<b>Transported minerals, low hematite content</b>	<b>1.53 – 0.0019 i</b>	<b>0.5</b>	<b>2.2</b>	<b>2.6</b>	<b>5.86 e-3</b>	<b>0.93</b>	<b>Dubovik, et al. 2002</b>

The set of 40 mixtures is meant to model all principally existing aerosol types and allow for some variability of the composition of each type. These set of mixtures has proven to provide a fit in the GOME spectra retrieval which is in many cases better than a 1% noise level.

Work to determine the success rate of the retrieval approach using this set of mixtures with a 14 month dataset is currently going on. In the old SYNAER aerosol model the (one) OPAC soot and mineral component was used.

Tab. 2: updated aerosol mixtures

No.		Name	Rel. hum. [%]	Vert. Prof. [km]	Component contributions to AOT550 [%]									
					WAS O	INSO	INSL	SSAM	SSCM	BISO	DISO	MITR	MILO	
1	21	Pure watersoluble	50/80	2	100									
2	22	Continental	50	2	95	5	5							
3	23				90	10	10							
4	24				85	15	15							
5	25	Maritime	50/80	2	30			70						
6	26				30			65	5					
7	27				15			85						
8	28				15			75	10					
9	29	Polluted watersoluble	50/80	2	90					10				
10	30				80					20				
11	31	Polluted Continental	50	2	80	10	10			10				
12	32				70	10	10			20				
13	33	Polluted Maritime	50/80	2	40			45	5	10				
14	34				30			40	10	20				
15	35	Desert Outbreak	50	2 - 4	25							75	75	
16	36			3 - 5	25								75	75
17	37			4 - 6	25								75	75
18	38	Biomass Burning	50/80	3	85					15				
19	39				70					30				
20	40				55					45				

WASO = watersoluble, INSO = insoluble, INSL = insoluble / low hematite, SSAM = sea salt accumul. mode, SSCM = sea salt coarse mode, BISO = biomass burning soot, DISO = diesel soot, MITR = mineral transported, MILO = mineral transported / low hematite; Mixture number N and mixture number N + 20: alternative humidity or mineral composition, respectively

#### 4. FIRST PRELIMINARY 14-MONTH CLIMATOLOGICAL DATASET

It is the ultimate goal of this work to produce (and in future update) a satellite based aerosol climatology for the observation area of the European geostationary Meteosat Second Generation satellite. For this purpose the SYNAER method will be implemented and work operationally with the sensors SCIAMACHY and AATSR onboard ENVISAT. A backup climatology production based on GOME and ATSR-2 products (both onboard ERS-2) of the year 1997/98 is shown here (using the old aerosol model). The final aim is to deliver 4 seasonal climatology datasets with a 1 degree horizontal grid. Due to cloud coverage and method inherent limitations a one year dataset was obtained as first product for a 5 degree grid.

For this first application of the SYNAER method data products of the period July 1997 through August 98 covering Europe/Africa were received through the ESA AO project SENECA (AO ID-106). Unfortunately, GOME measures "small" pixels of 80x40 km<sup>2</sup> only for

3 days every month, and 320x40 km<sup>2</sup> pixels throughout the rest of the month. For producing the aerosol climatology only "small" pixels are meaningful in correspondence with horizontal variability scales of the aerosol loading and are thus exploited. With this small pixel mode GOME covers a swath width of 240 km with only 3 pixels in one scan line (SCIAMACHY will deliver pixels of 60x30 km<sup>2</sup> for a swath of 960 km, i. e. 16 pixels in one line). Although this means a severe limitation to the available data base, it opens the opportunity to test the methodology with a one year dataset. Detailed handling of quality information from the retrieval process such as fit error, GOME-ATSR-2 cross-calibration deviation, spectral noise, surface elevation, solar elevation angle, etc. was optimized. Depending on the cloud coverage and method inherent limitations (surface brightness must not exceed 8% over land and 1.5 % over ocean, GOME pixel cloud fraction must not exceed 35%, differentiation of aerosol types is only reliable for optical thickness at 550 nm larger than 0.1, the ambiguity test rejects pixels with a fit error less than 0.005) extracting the seasonality was hardly possible with this first dataset. Given the limited temporal coverage of 3 observations per month (which

is further reduced due to cloudiness and technical errors) it was assured during averaging of the climatology values that a minimum of 2 orbits (i.e. 2 different times of observations) and of 5 pixels contributed to the box average (mean values for grid boxes with result values are 6 orbits and 41 pixels). To

obtain a significant result a 14 month dataset (average and standard deviation optical thickness, number of pixels contributing to one grid box, differentiation of optical thickness into basic components) using a total number of 18914 GOME pixels was produced as first product for a 5 degree grid.

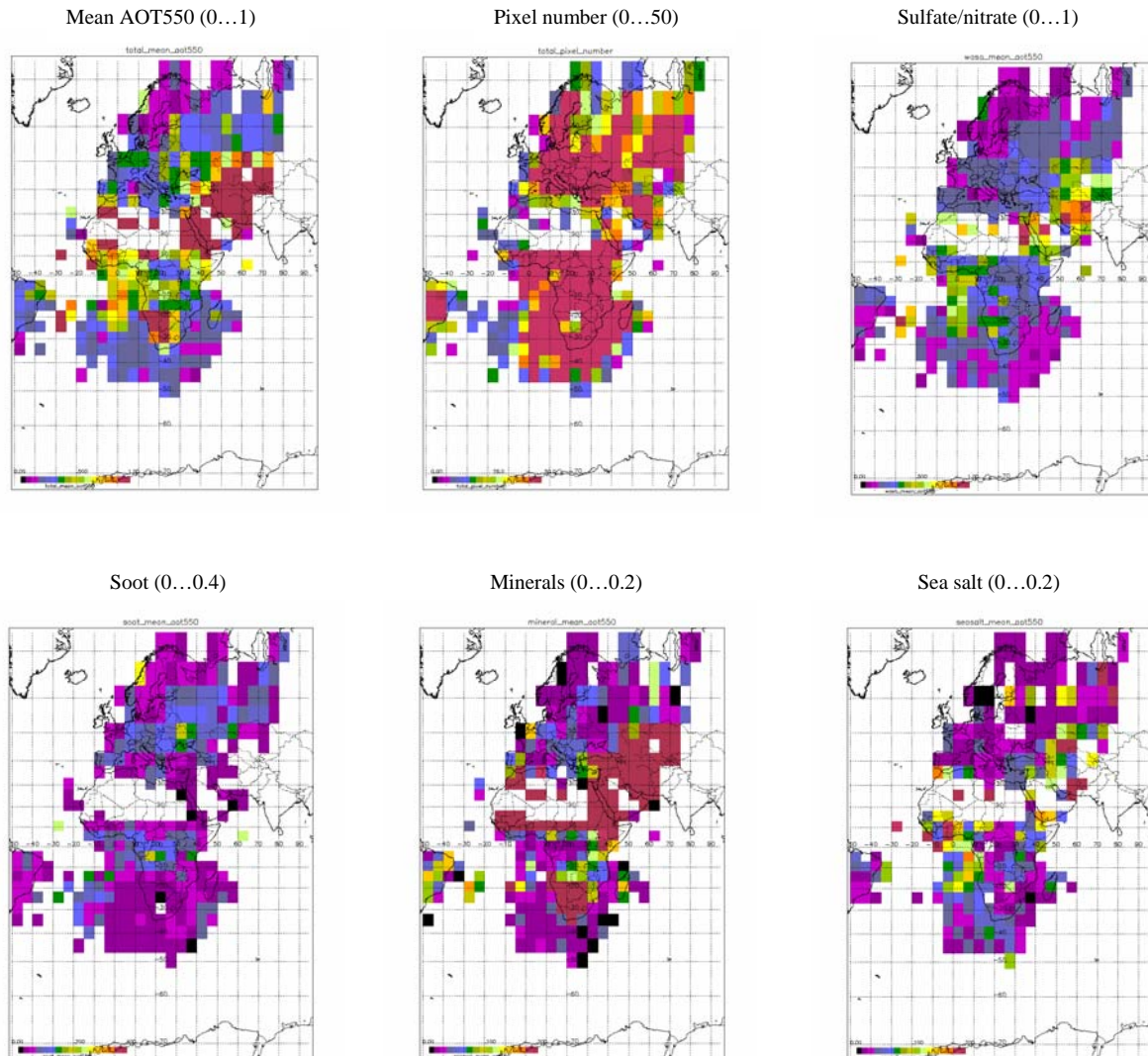


Fig. 1: climatology of aerosol components

Fig. 1 shows the climatology dataset based on the 14 months of GOME/ATSR-2 data on a 5 degree grid. The low number of observations in some boxes (particularly in the South-West Atlantic and at the edges of the dataset) leads to some unrealistic mean concentrations values which are determined by single episodes. This will be enhanced by using the daily ENVISAT observations. However, despite of the small data base the aerosol distribution is rather smooth with a mean aerosol optical thickness of 0.26 and some interesting features can already be observed: Largest optical thickness values above 1 occur over / near the

desert areas, whereas the Scandinavian area or oceanic zones far off from the continents show lowest values. Also biomass burning plumes from South America and Central Southern Africa are indicated over the Atlantic.

Fig. 1 shows also the component-wise mean aerosol optical thickness maps: Sulfate/nitrate aerosols which are included in all modelled aerosol types as background contribution show even clearer the unpolluted oceanic and Northern areas. Soot occurs most prominently over industrialized / densely populated areas in Central/Eastern Europe as well as

over biomass burning source areas in Central/Southern Africa and in their plumes over the ocean. Maximum 14 month 5 degree box average values are near 0.4. Few grid boxes show exceptionally high values which are due to a low number of contributing pixels. Dust is dominant in and around desertic areas (Sahara, Namib, Near East) and occurs with smaller amount in continental and oceanic aerosol. Sea salt occurs with low values also at inland locations which indicates to the limits of separating this component with low impact on the total optical thickness, but it should be noted that its maximum occurs over the Southern Atlantic. Some areas are not covered by the basic dataset because of either too high cloudiness or too bright surface. Due to miss-interpretation because of the simplified dust vertical structure high soot and sea salt values occur over deserts to some extent. Furthermore, an error in the land sea mask was found in the South-East-Atlantic which leads to some increased concentration levels over the ocean. Finally, new ground-based measurements revealed limitations in the soot and dust representation in the SYNAER aerosol model (see section 3), which are also a potential reason of wrong aerosol type detection.

## 5. FIRST ENVISAT RESULTS

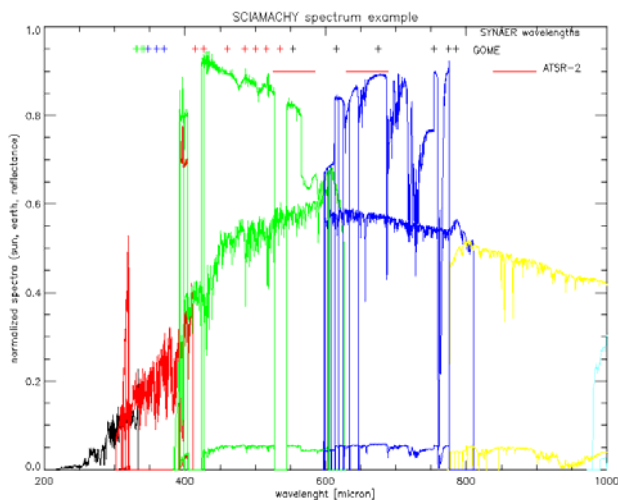


Fig. 2: SCIAMACHY normalized spectra (reflectance, sun radiation, earth radiation; from top to bottom) from orbit 13044 (28 August 2004) over Northern Russia (64.6N/13.1E); only the UV-visible and near-infrared range is plotted; + indicates SYNAER retrieval wavelengths \_\_\_ show the ATSR-2 bands.

The SYNAER method was prepared for ENVISAT AATSR and SCIAMACHY data by extracting exactly the same information from both sensors which was

available from GOME and ATSR-2. Namely, the orbit-wise data format of AATSR was split into frames of 512x512 pixels and the wider spectral coverage of SCIAMACHY was reduced to the GOME wavelength range. However, due to the mission concept (e.g. SCIAMACHY is permanently switching between nadir and limb observations; shortest integration time associated with highest spatial resolution is limited to parts of the entire spectrum, so-called clusters) not exactly the same information content is available. This can be seen in fig. 2, where gaps in the SCIAMACHY example spectrum are visible. Therefore a fine tuning of retrieval wavelengths (indicated in fig. 2 as + at the top) is necessary.

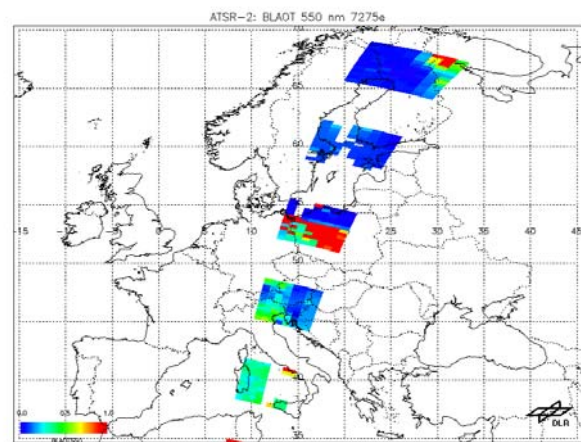


Fig. 3: example ENVISAT aerosol optical thickness map (orbit 7275 of 22 July 2003)

Furthermore the SCIAMACHY-AATSR correlation accuracy of radiometric calibration and geolocation as well as instrument specific properties such as the "spatial aliasing" (non-negligible read-out time as compared to integration time leading to a moving field of view during the spectral scan) need to be analyzed. As a first attempt correction factors based on an intercomparison to MERIS in the visible bands after [15]. (which agree to the ones by [16]) have been used. Calculation of SCIAMACHY reflectances required an accurate spectral sun-to-earth spectra coregistration. Fig. 3 shows a first preliminary ENVISAT aerosol optical thickness result for orbit 7275 of 22 July 2003 (synthetic SCIAMACHY pixels; fixed aerosol type: 100% watersoluble). The much better spatial resolution and coverage (as compared to ERS-2) and the gaps due to the limb-nadir sequence are clearly visible. In the center the low optical thickness values in the Alpine bow are evident. A rough intercomparison of the results with AERONET sun photometer measurements at Ispra and Venise shows qualitatively good agreement

better than 0.1. Very high optical thickness values (in red) are probably due to erroneous cloud detection. Here, the operational AATSR cloud flag was used; one future remaining task is the utilization of the APOLLO scheme as with ERS-2, which provides 1km cloud fraction results.

## 6. SUMMARY AND CONCLUSION

The potential of the SYNAER method for separating basic aerosol components has been demonstrated with a limited 14 month database. The method will be optimized (especially with regard to absorbing aerosol components) and used for the production of an aerosol component climatology from space. For this purpose SYNAER is adapted to ENVISAT SCIAMACHY and AATSR to produce daily global aerosol maps with 480 km swath and 60x30 km<sup>2</sup> pixel size. Furthermore, simplifications in the current method setup, which are responsible for miss-selection of the aerosol type in certain conditions are currently revised to lead to a more accurate optical thickness retrieval and differentiation of the aerosol type. Then, a reprocessing of the entire GOME/ATSR-2 background 14 month climatology dataset and operational ENVISAT aerosol optical thickness and type processing (from 2005) will be conducted (NRT: over Europe, OL: global), which takes these improvements into account. Based on the thus becoming available large data amount systematic validation and cross-comparison with ERS-2 observations in 2002/2003 will be conducted. SYNAER delivers relevant information which is required for aerosol data assimilation (national DFG project AERO-SAM) in air quality forecasting, for monitoring the different components of the atmospheric aerosol loading (EUMTSAT climate SAF), and for the provision of accurate solar radiation to solar energy technologies (EOMD project ENVISOLAR). A long-term application scenario with similar sensors onboard ERS-2, ENVISAT and METOP spanning the period 1995 to 2020 is envisaged at DLR-DFD

## 7. ACKNOWLEDGEMENT

The results presented in this paper were achieved within the ESA-AO projects PAGODA (AO2.D107-1) and PAGODA-2 (AO3.218) and ESA-ENVISAT-AO project SENECA (AO ID-106) through which the input data (GOME and ATSR-2; SCIAMACHY and AATSR) were acquired as well as within the EU FP5 project HELIOSAT-3 (NNE5-2000-00413FP5).

## 8. BIBLIOGRAPHIC REFERENCES

1. Holzer-Popp, T., M. Schroedter, and G., Gesell, Retrieving aerosol optical depth and type in the boundary layer over land and ocean from simultaneous GOME spectrometer and ATSR-2 radiometer measurements, 1, Method description *J. Geophys. Res.*, 107, D21, pp. AAC16-1 – AAC16-17, 2002a
2. Holzer-Popp, T., M. Schroedter, and G., Gesell, Retrieving aerosol optical depth and type in the boundary layer over land and ocean from simultaneous GOME spectrometer and ATSR-2 radiometer measurements, 2, Case study application and validation, *J. Geophys. Res.*, 107, D24, pp. AAC10-1 – AAC10-8, 2002b
3. Pope III, C. A., Burnett, R. T., Thun, M. J., Calle, E. E., Krewski, D., Ito, K., Thurston, G. D., *J. Am. Med. Ass.*, 287, 1132-1141, 2002
4. Stedman J. R., The predicted number of air pollution related deaths in the UK during the August 2003 heatwave, *Atmos. Env.*, 38, 1087 – 1090, 2004
5. Pohl, O., News scan: Disease Dustup, *Sci. Am.*, 10-11, July 2003
6. Prospero, J. M., Ginoux, P., Torres, O., Nicholson, S., Gill, T., Environmental characterization of global sources of atmospheric soil dust identified with the NIMBUS 7 Total Ozone Mapping Spectrometer (TOMS) absorbing aerosol product, *Rev. of Geophys.*, 10.1029/2000RG000095, 2002
7. Petzold A., Lauer A., Kaminski U., Klimawirksamkeit des schwarzen Kohlenstoffs verlangt genauere Messverfahren: Erste Ergebnisse mit einem neuartigen Mehrwinkelabsorptionsphotometer, GAW Brief des Deutschen Wetterdienstes, 03/04, DWD Hohenpeißenberg, 2004
8. Kaufman, Y. J., Tanre, D., Boucher, O., A satellite view of aerosols in the climate system, *Nature*, 419, 215 – 223, 2002
9. Hess, M., Köpke, P., Schult, I., Optical Properties of Aerosols and Clouds: The Software package OPAC, *Bulletin of the Americal Meteorological Society*, 79, pp. 831-844, 1998
10. Kriebel K. T., Gesell G., Kästner M., Mannstein H., The cloud analysis tool APOLLO: Improvements and Validation, *Int. J. Rem. Sens.*, 24, 2389-2408, 2003
11. Myhre, G., Stordal, F., Johnsrud, M., Diner, D. J., Geogdzhayev, I. V., Haywood, J. M., Holben, B., Holzer-Popp, T., Ignatov, A., Kahn, R., Kaufman, Y. J., Loeb, N., Martonchik, J., Mishchenko, M. I., Nalli, N. R., Remer, L. A., Schroedter-Homscheidt, M., Tanre, D., Torres, O., Wang, M., Intercomparison of satellite retrieved aerosol optical depth over ocean during the period September 1997 to December 2000, submitted, *J. Geophys. Res.*, 2004
12. Dubovik, O., B. Holben, T.F. Eck, A. Smirnov, Y.J. Kaufman, M.D. King, D. Tanre, I. Slutsker,

Variability of Absorption and optical Properties of Key Aerosol Types Observed in Worldwide Locations, *J. Atm. Sciences*, Vol 59, 590 – 608, 2002

13. Schnaiter, M., H.Horvath, O. Möhler, K.-H. Naumann, H. Saathoff, O.W. Schöck, UV-VIS-NIR spectral optical properties of soot and soot-containing aerosols, *J. Aerosol Sci.*, 34, 1421-1444, 2003

14. Sinyuk, A., O. Torres, O. Dubovik, Combined use of satellite and surface observations to infer the imaginary part of refractive index of Saharan dust, *Geophys. Res. Lett.*, Vol 30, No 2, 1081, doi 10.1029/2002GL016189, 2003

15. Von Hoyningen-Huene, W., Kokhanovsky, A. A., Wuttke, M. W., Buchwitz, M., Noel, S., Gerilowski, K., Burrows, J. P., Latter, B., Siddans, R., Kerridge, B. J., Validation of SCIAMACHY Top-of-Atmosphere Reflectance for Aerosol Remote Sensing Using MERIS L1 Data, accepted for publication in *Atmos. Chem. Phys.*, 2004

16. De Graaf, M., Stammes, P., Verification of SCIAMACHY's Absorbing Aerosol Index product, Proc. of ENVISAT Validation Workshop, ESRIN, Frascati, 9.-13. Dec. 2003, ESA SP-531, August 2003, 2003